

Examining Relationships Between the Vertical Structure of Deep Convection and Upper Tropospheric Humidity Using AIRS



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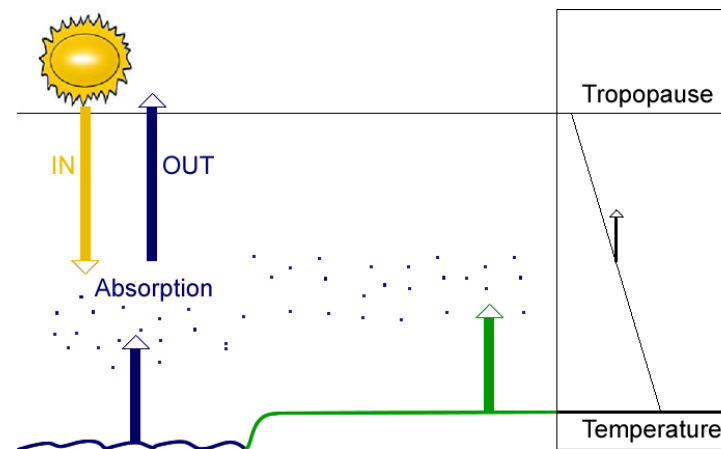
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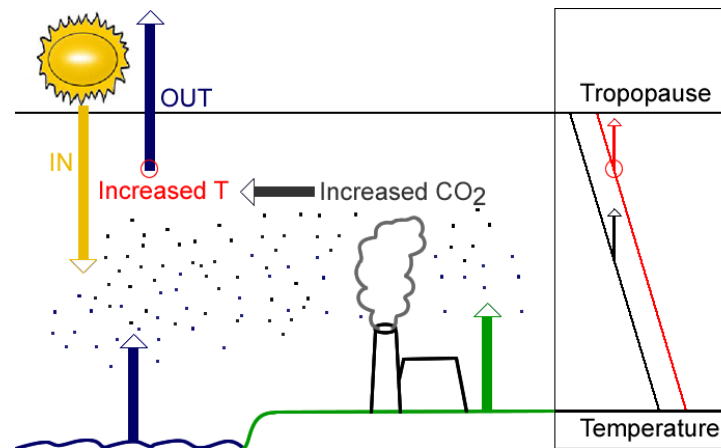
- Introduction
- Data & Method
- Preliminary Results
- Future Work



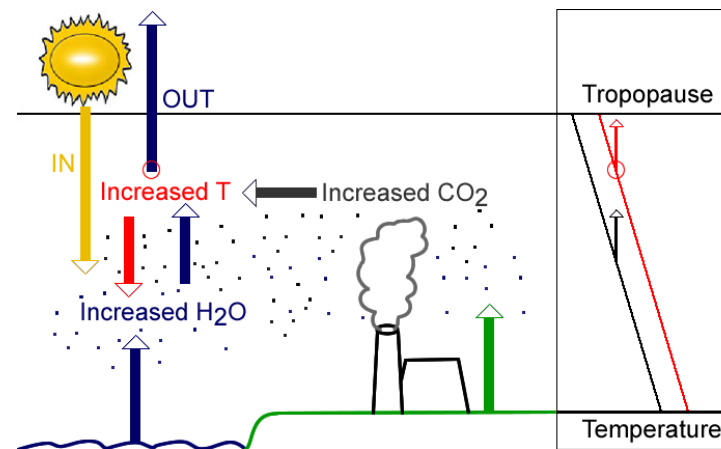
- Water vapor: the dominant greenhouse gas
 - ▷ Continuum absorption in IR
 - ▷ Abundance in atmosphere
- Atmospheric capacity for water vapor increases with increasing temperature \Rightarrow expect feedback to temperature changes



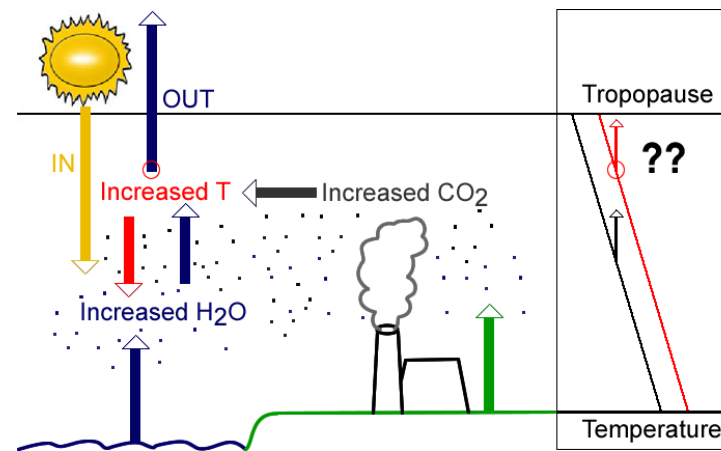
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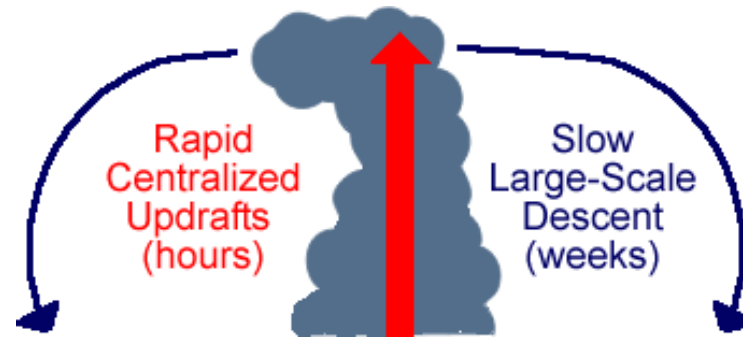


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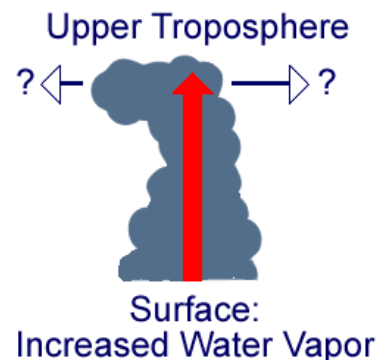
- Strength of feedback remains uncertain: estimates range from zero feedback to constant RH ($\sim 170\%$), or more!

- Climate models: 35% of total radiative water vapor feedback from tropical UTH (100-500 mb)
- Cold temperatures in tropical, subtropical UT mean that a small change can have a large effect
- Conceptual model of tropical upper tropospheric water vapor:
 - ▷ Source: rapid, highly localized convection
 - ▷ Sink: slow, large scale descent



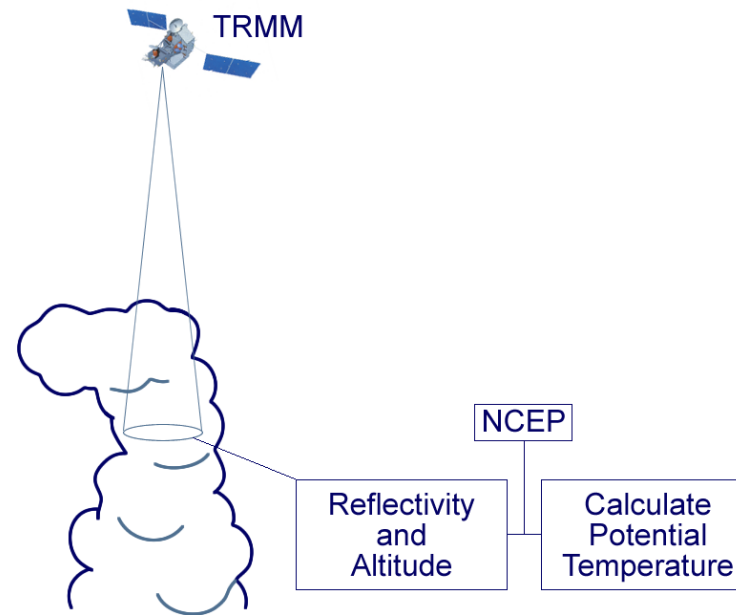
- Water vapor distribution largely controlled by distribution of convection

- Convection can both hydrate and dehydrate the UT
 - ▷ Retention and evaporation of droplets \Rightarrow moistening
 - ▷ Vapor condenses onto droplets and precipitates \Rightarrow dehydration
 - ▷ Detrainment into already saturated air, drops fall out \Rightarrow no change
- Current climate models: moisture detrainment controlled by temperature (altitude) of detraining layer
- Other influences: cloud/precip microphysics, mesoscale downdrafts
- Strength of modeled water vapor feedback highly dependent on detrainment scheme

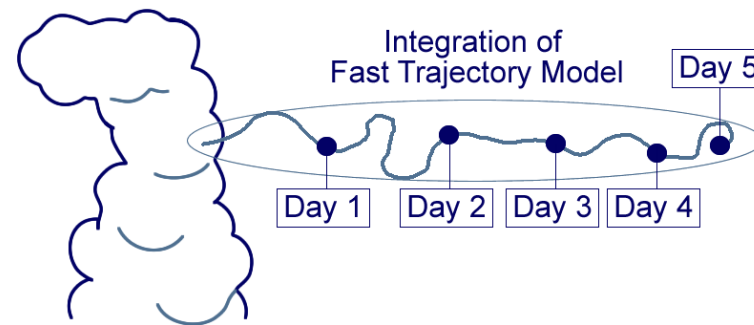


- Previous studies of convective detrainment in the UT:
 - ▷ *in situ*: highly localized observations of short term evolution
 - ▷ Models: larger scale, longer term but necessarily simplified physics
 - ▷ Satellites: vertical structure unknown, water vapor observations sparse
- Recent satellite technology provides unprecedented opportunities
 - ▷ TRMM Precipitation Radar: vertical characterization of convective systems
 - ▷ AIRS: high vertical resolution global coverage of water vapor into the upper troposphere
 - ▷ MODIS: Ice particle sizes at cloud top
- Link these observations by a transport scheme
- Preliminary proof of concept study:
 - ▷ Detrainment altitude
 - ▷ Cloud/precip microphysics
 - ▷ Role of ice in UT water vapor feedback

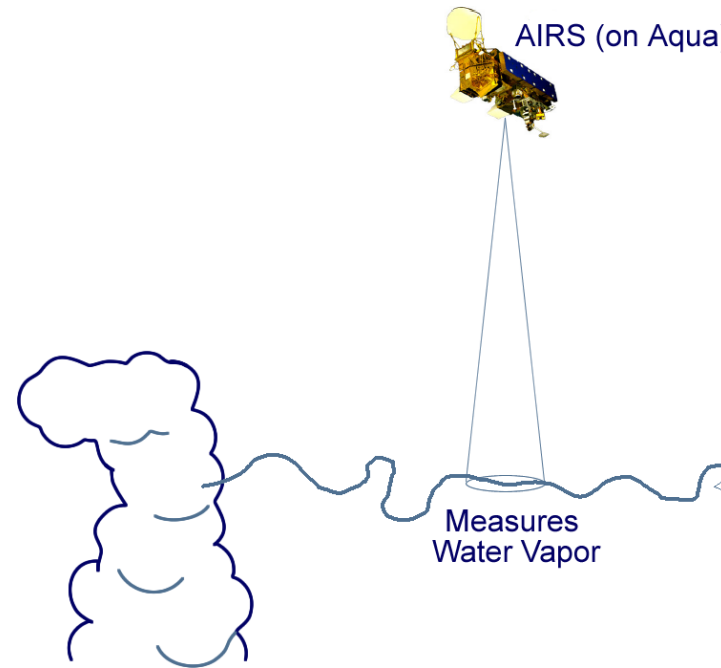
- TRMM Precipitation Radar
 - ▷ 2A25 Volumetric Radar Reflectivities
 - Echo from water and ice droplets within a volume
 - Higher reflectivities = larger droplets or higher concentrations
 - Measure of convective intensity
 - ▷ Reliable for convective systems larger than footprint (4.3 to 5 km)
- AIRS
 - ▷ Combination of IR and microwave instruments
 - ▷ Rapid global coverage ($\sim 2\times$ per day)
 - ▷ Horizontal resolution ~ 40 km at nadir; vertical resolution ~ 2 km.
 - ▷ Slight dry bias in upper troposphere relative to ECMWF
- MODIS
 - ▷ Cloud ice particle effective radius derived from visible and infrared radiances
 - ▷ Along track or daily $1^\circ \times 1^\circ$ gridded product



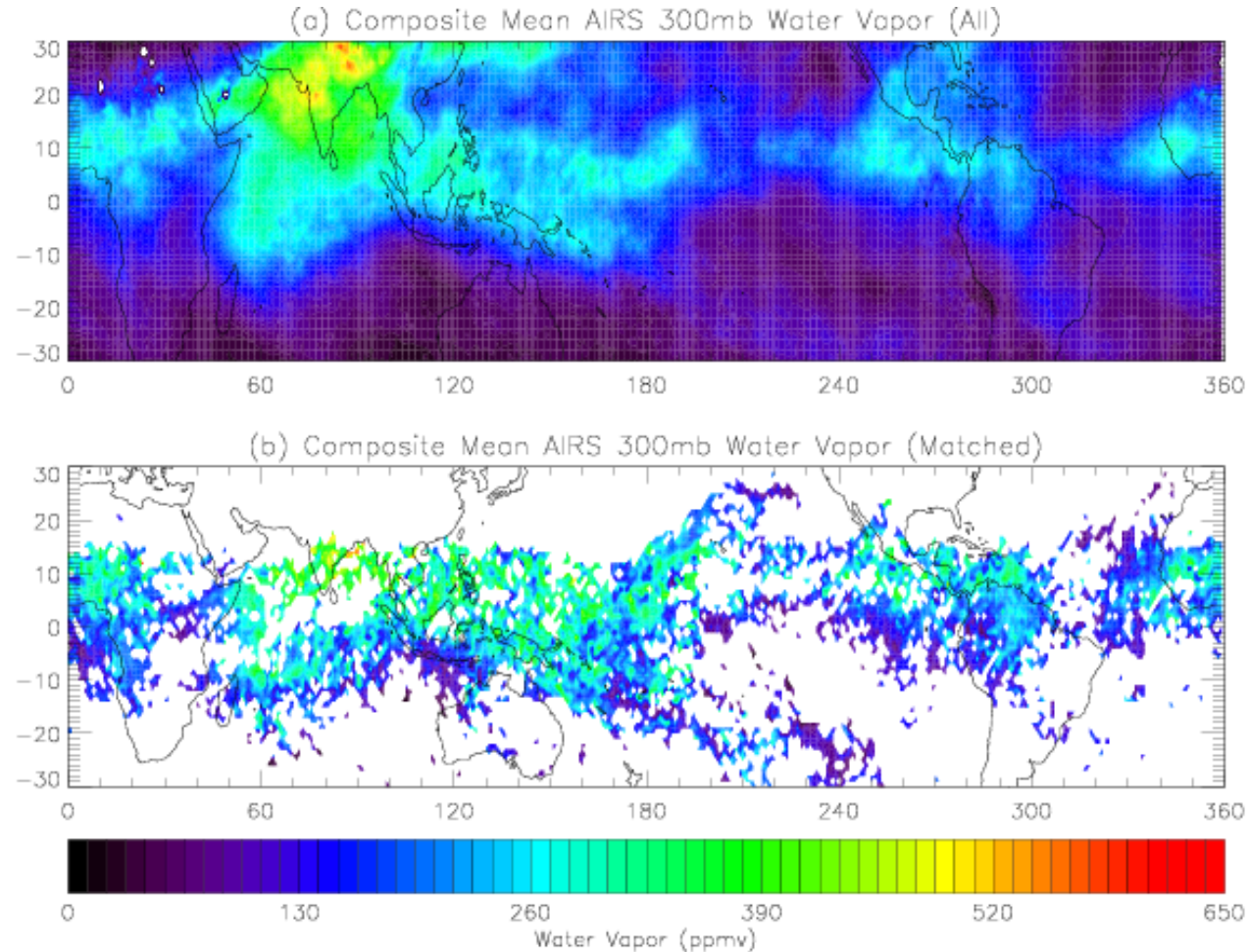
- Scan TRMM observations for:
 - ▷ Deep convection (altitude ≥ 10 km)
 - ▷ TRMM PR $Z \geq 20$ dBZ (noise threshold ~ 17 dBZ)
- Calculate potential temperature from NCEP geopotential heights, assume TRMM altitude \equiv NCEP geopotential height, and interpolate
- Store MODIS mean cloud ice effective radius for associated gridbox



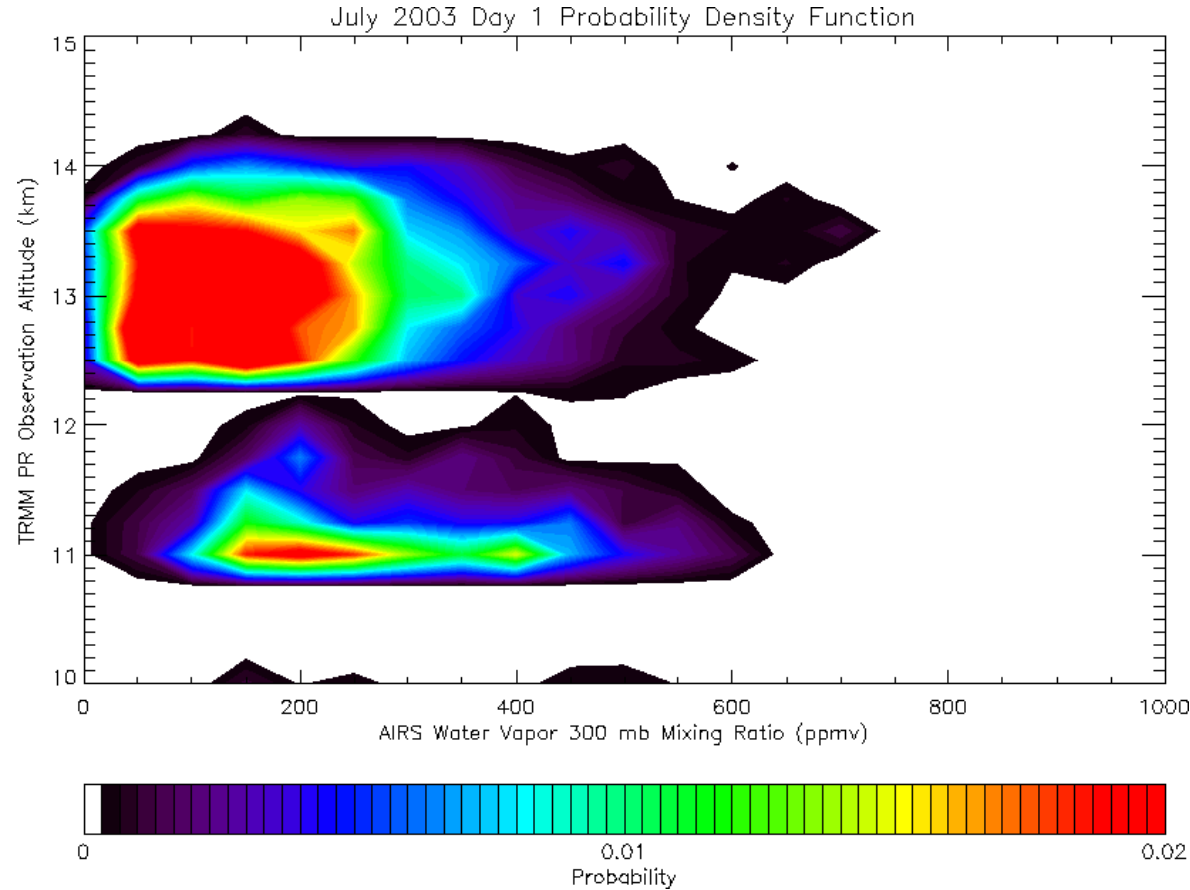
- Fast Trajectory Model - ftraj (M. Schoeberl)
 - ▷ Five day forward trajectory with timestep = 0.02 days (~ 30 minutes)
 - ▷ UKMO winds (Updated daily at 12 UTC, 2.5° lat \times 3.75° lon)
 - ▷ Diabatic heating rates derived from UKMO using a radiative transfer scheme
- Position stored at each timestep



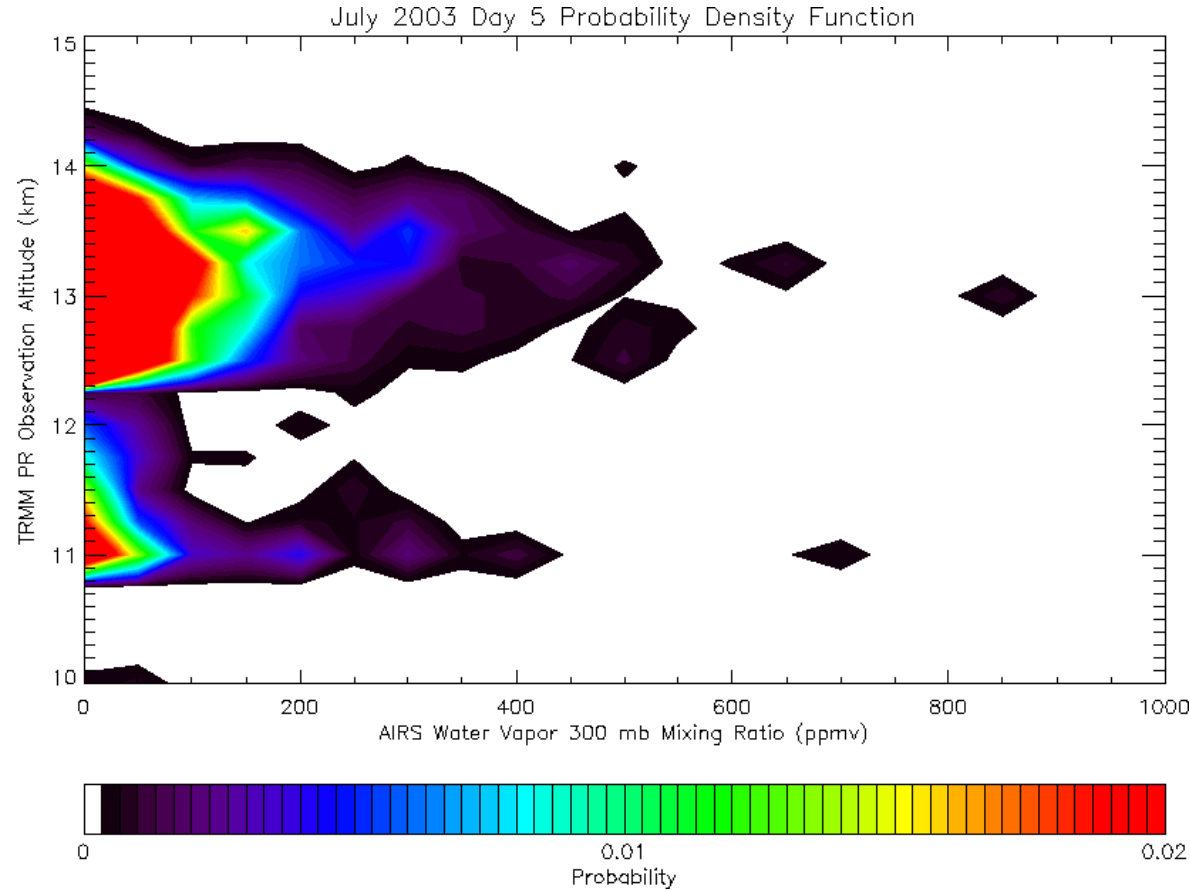
- Search for AIRS observations close in space and time to trajectory point
 - ▷ $1^\circ \times 1^\circ$ box & 30 minutes following trajectory passage
 - ▷ Include unvalidated overland measurements
- If multiple locations, use mean humidity
- Linearly interpolate from AIRS standard pressure levels



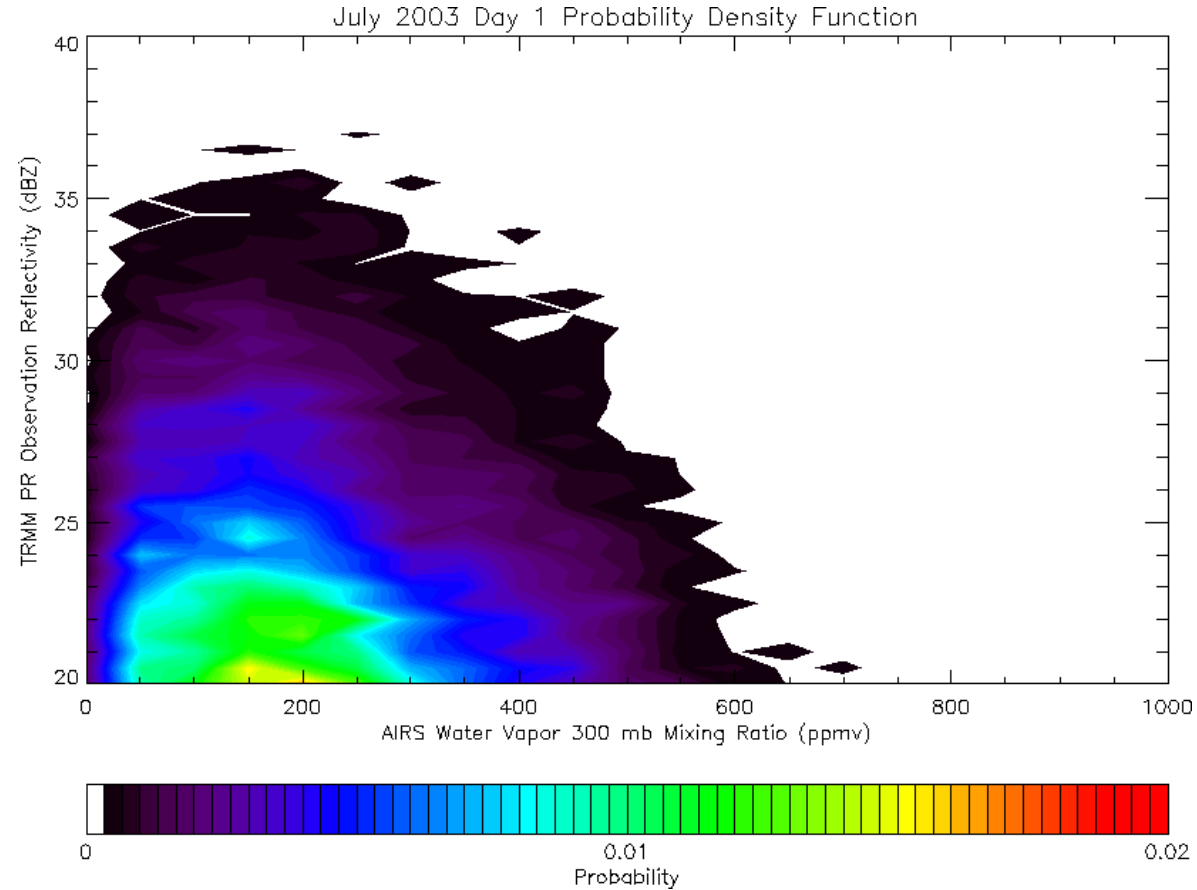
- ▷ Many of the maxima are influenced by convective events observed in TRMM
- ▷ Consistent with conceptual model - bolsters confidence in the method



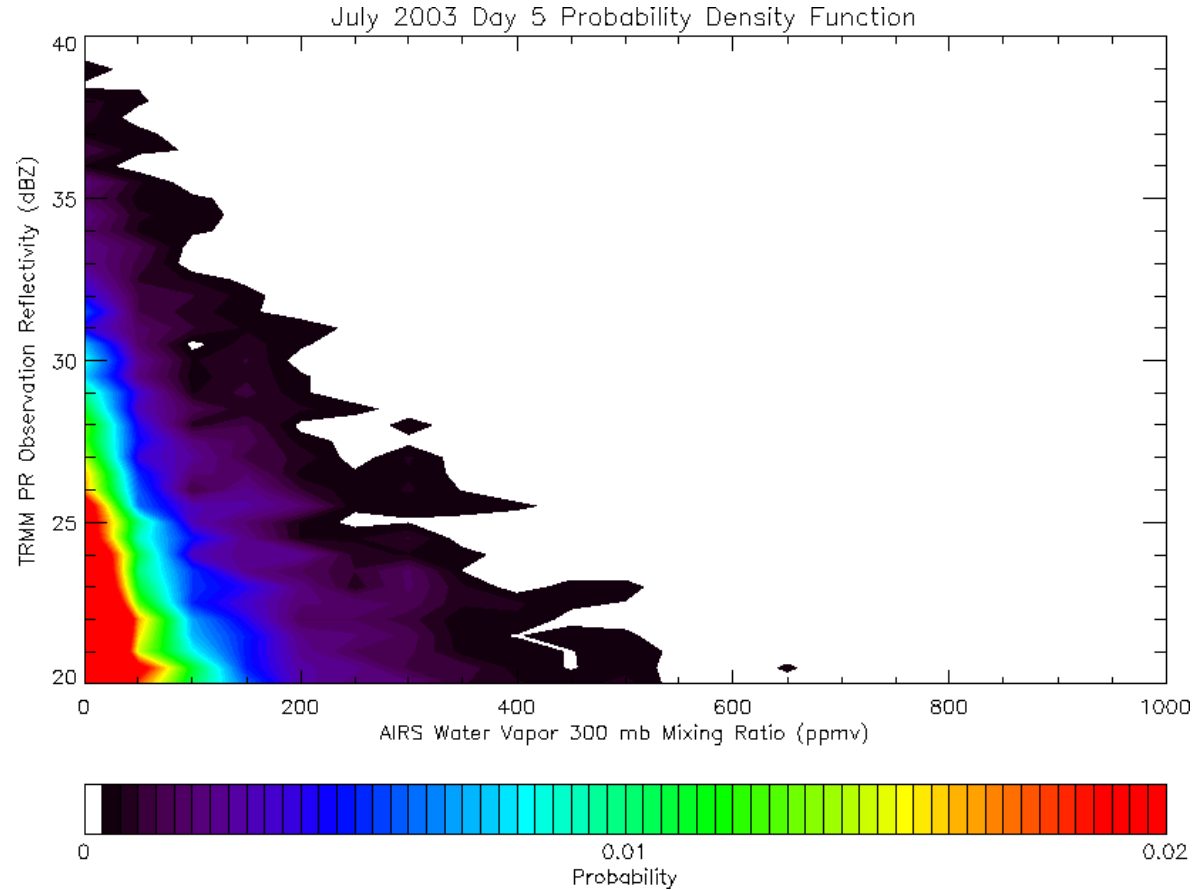
- ▷ Apparent bimodal outflow distribution: 11-12 km, 12.5-14 km
- ▷ Outflow altitude looks too high! Likely due to estimation of θ



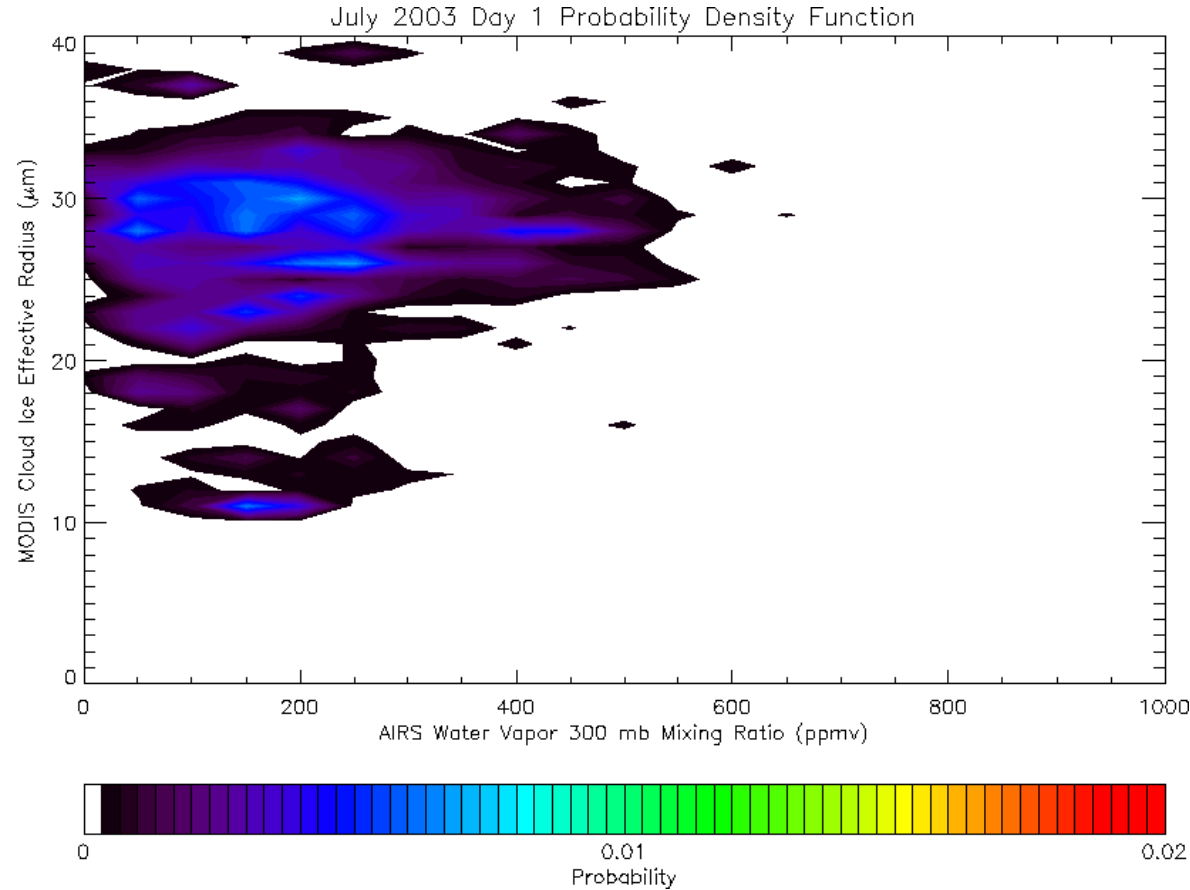
- ▷ Apparent bimodal outflow distribution: 11-12 km, 12.5-14 km
- ▷ Outflow altitude looks too high! Likely due to estimation of θ
- ▷ Higher altitudes may dehydrate more slowly; gap blurs



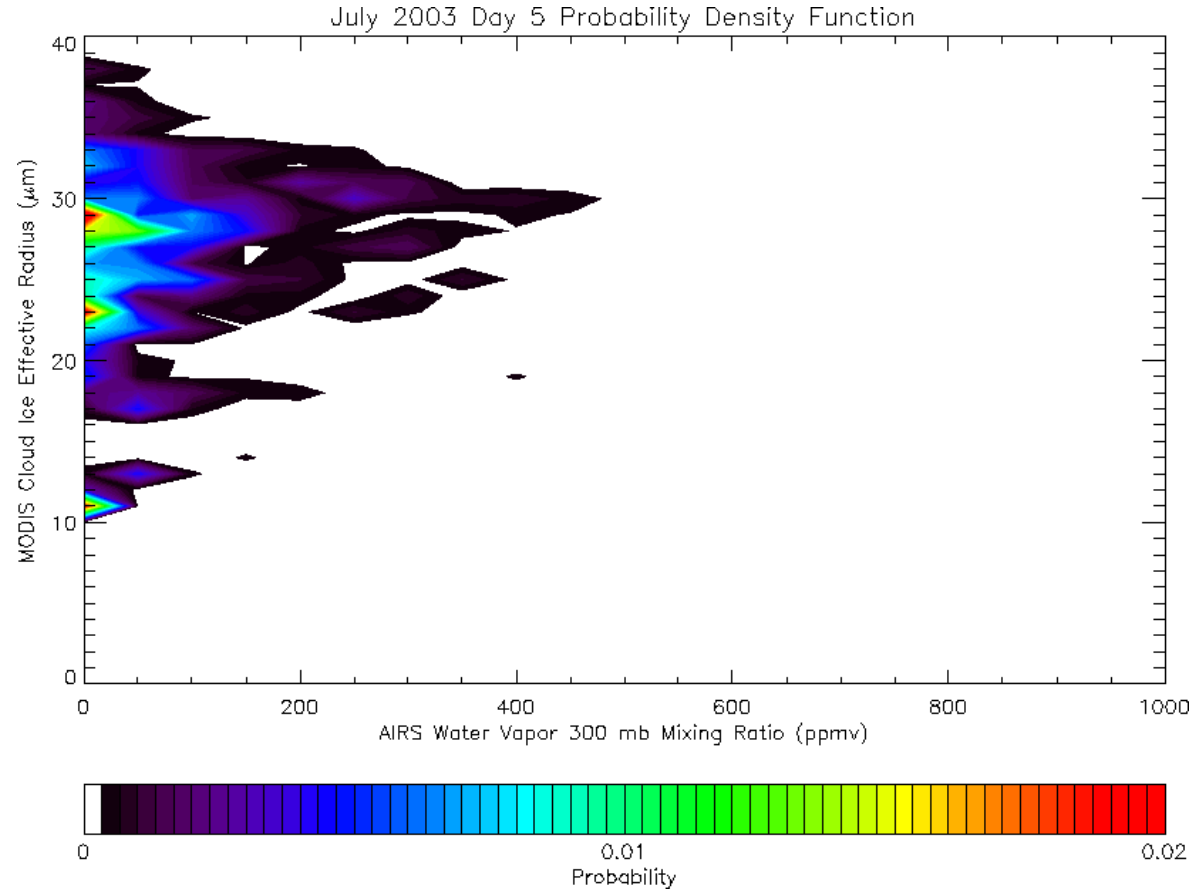
- Stronger convection seems to detrain drier air



- ▷ Stronger convection seems to detrain drier air
- ▷ Detrainment from higher reflectivities appears to dehydrate more quickly
- ▷ Stronger convection \Rightarrow higher precip efficiency \Rightarrow drier air downstream



- ▷ Main cluster between 20 and 35 μm
- ▷ Smaller effective radius/lower humidity due to higher detrainment altitude?



- ▷ Main cluster between 20 and 35 μm
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- ▷ Evaluate gridded vs. along-track

- Preliminary results indicate:
 - ▷ Detrainment at higher altitudes may dehydrate more slowly
 - ▷ Bimodal distribution of detrainment - continental vs. maritime convection?
 - ▷ Larger reflectivities may dehydrate more quickly
- Estimation of potential temperature a major weakness
- Need to evaluate MODIS results, particularly level 3 vs. level 2
- Otherwise, the method and data used in this preliminary study show significant potential for use in broader and longer term studies
 - ▷ Develop method to check for cirrus along track (ISCCP DX)
 - ▷ Investigate regional/seasonal variability over 2 years
 - ▷ Case studies: bin trajectories by system; match with aircraft studies
 - ▷ “Train” mixing parameterization along trajectory by tracking individual trajectories
 - ▷ Evaluate role of boundary layer aerosols (e.g., biomass burning)